

1 1 **A spatially hybrid hydrological modeling approach using**
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4 2 **subbasin-specific model structures**

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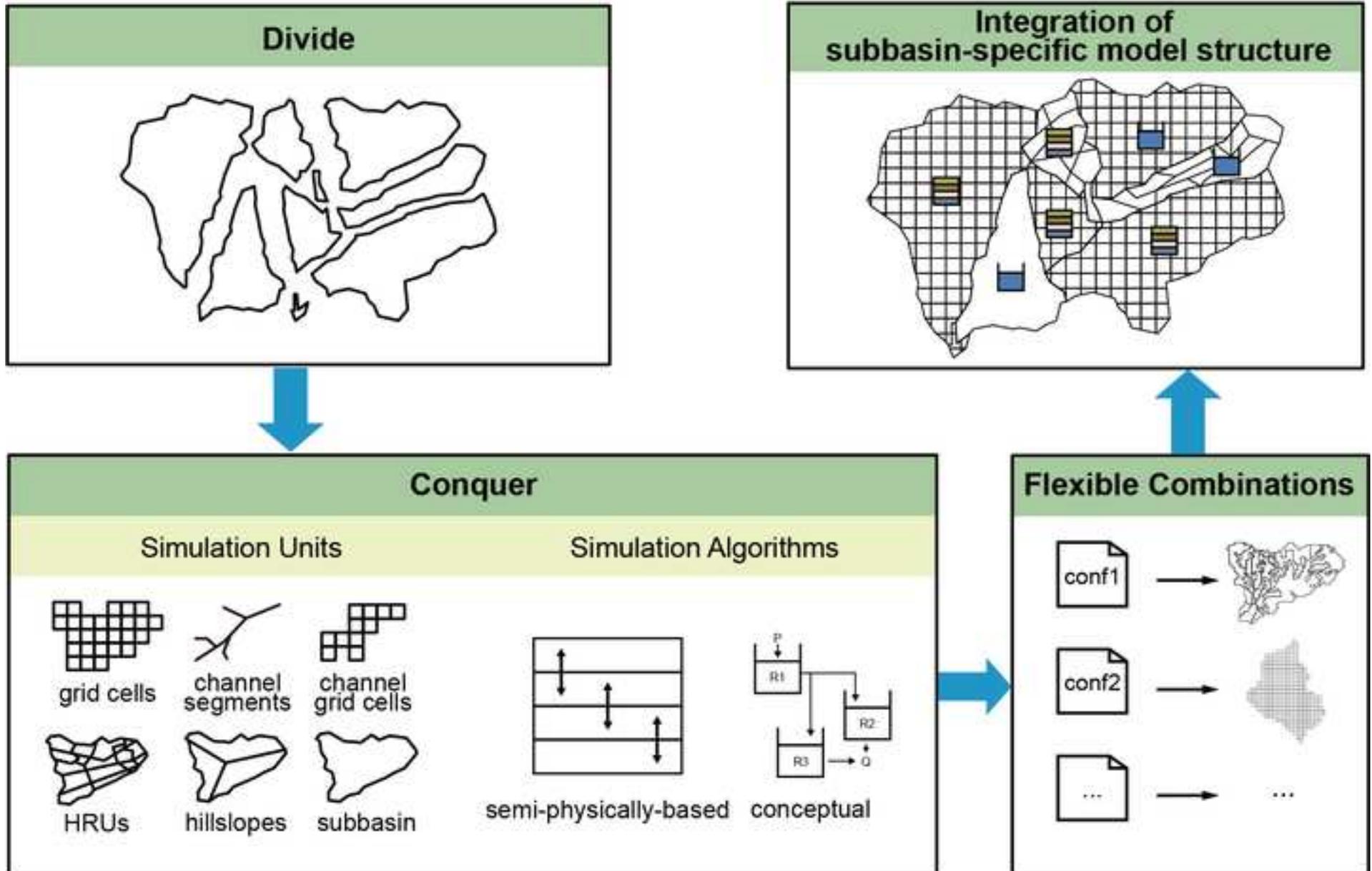
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Highlights

- An approach to supporting spatially hybrid hydrological model structures is proposed.
- Diverse subbasin-specific model structures are integrated for watershed-scale simulation.
- Model structure flexibly combines compatible spatial units and simulation algorithms.
- Spatial units of subbasin, hillslope, HRU, and cells support conceptual or semi-physical algorithms.
- Spatially hybrid model from the proposed approach improves general performance.

Spatially Hybrid Hydrological Modeling Approach



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32 **Abstract:**

33 Hydrological models that typically adopt spatially consistent model structures
34 are often unfit for complex application contexts with significant spatial heterogeneity.
35 Although existing hydrological modeling frameworks support spatially varying
36 lumped or semi-distributed conceptual model structures, they face a significant
37 challenge in integrating distributed physically-based model structures. This paper
38 proposes a new spatially hybrid hydrological modeling approach, which flexibly
39 combines compatible spatial units and simulation algorithms to construct different
40 model structures for individual subbasins within interest watershed, and then
41 coherently integrate them through channel routing to perform watershed-scale
42 simulation. Implemented within the Spatially Explicit Integrated Modeling System,
43 this approach was evaluated in a medium-sized watershed with two distinct model
44 structures: a lumped conceptual model structure for gently sloping subbasins and a
45 fully-distributed semi-physically-based model structure for mountainous subbasins.
46 Comparative experiments demonstrated that the hybrid model combines the strengths
47 of both, improving general performance. The proposed approach enhances modeling
48 flexibility towards knowledge-driven intelligent modeling in spatially heterogeneous
49 environments.

50 **Keywords:** Watershed simulation; hydrological modeling framework; spatially hybrid
51 modeling; model structure; spatial heterogeneity; SEIMS

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1 Introduction

Hydrological models serve as effective tools for simulating and enhancing our understanding of complex hydrological processes in real-world watersheds. A variety of hydrological models have been developed, each suited to different application contexts depending on multiple factors such as watershed characteristics and available data (Beven, 2000; Gharari et al., 2021; Wagener et al., 2001). Building an application-specific model that is both hydrologically reasonable and practically executable stands as the primary and critical step for hydrological modeling and applications such as scenario analysis of watershed management (Butts et al., 2004; David et al., 2022; Pilz et al., 2020; Qin et al., 2025; van Esse et al., 2013).

The model structure of a hydrological model can be defined by its spatial discretization scheme and the representation of hydrological processes (Chow, 1988; Milad et al., 2012; Hrachowitz and Clark, 2017). The spatial discretization scheme refers to the arrangement of one or more types of spatial units (also known as simulation units) within the watershed for hydrological simulation (Dehotin and Braud, 2008). Lumped spatial units typically ignore or broadly generalize the spatial heterogeneity of watershed characteristics, including either coarse-scale units (such as a subbasin or even an entire watershed), or units defined by dominant combinations of selected geographic attributes (Arnold et al., 2010). In contrast, semi- or fully-distributed spatial units can fully cover the entire geographic spatial domain of the watershed at varying levels of details, such as spatially explicit HRUs (hydrological

1 88 response units; Teshager et al., 2016) and grid cells (Wigmosta et al., 1994). The
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4 89 representation of hydrological processes denotes the simulation algorithms used for
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6 90 hydrological phenomena (e.g., infiltration and surface flow routing). Conceptual
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9 91 representations simplify hydrological processes with a few lumped components,
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12 92 typically using conceptual reservoirs (or buckets) (Fenicia et al., 2011; Knoben et al.,
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15 93 2019). In contrast, fully-physically-based representations apply fundamental physical
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18 94 laws (e.g., mass, momentum, and energy conservation) by solving coupled partial
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21 95 differential equations (Abbott et al., 1986; Shu et al., 2020). To balance physical
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24 96 realism with computational feasibility, semi-physically-based representations employ
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27 97 simplified physical equations for individual hydrological processes that are generally
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30 98 simulated independently (Arnold et al., 1998; Tague and Band, 2004; Tang et al.,
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33 99 2019).

100 Existing methods of determining model structures can be categorized as
101 spatially consistent method and spatially varying method. The spatially consistent
102 model structure means the types of simulation units and simulation algorithms for
103 interested hydrological processes are the same for the whole watershed. Most research
104 on directly choosing a single existing hydrological model with a fixed model structure
105 falls into this category. A more flexible choice is utilizing modular hydrological
106 modeling frameworks, which emphasize various aspects of flexibility in customizing
107 application-specific model structures, such as extensible types of simulation units and
108 associated simulation algorithms (Kneis, 2015), and alternative simulation algorithms

1 109 for individual hydrological processes (Clark et al., 2015; Craig, 2020; Knoben et al.,
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4 110 2019; Zhu et al., 2019). However, their default design or most common application
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6 111 remains on customizing and applying a spatially consistent model structure. This
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9 112 method inherently assumes that the spatial variability of watershed characteristics can
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12 113 be adequately captured by the required input data and parameters. However, using a
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15 114 fixed model structure may fail to effectively accommodate diverse and complex
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18 115 application contexts exhibiting significant spatial heterogeneity (Ley et al., 2016;
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21 116 Savenije, 2009; Sui and Turner, 2021).

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23 117 Recognizing the limitation of spatially consistent model structures, researchers
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26 118 have explored constructing spatially varying model structures within watersheds in
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29 119 two principal approaches. The first involves externally integrating multiple individual
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32 120 models, either manually (Liu et al., 2020; Li et al., 2021; Wang et al., 2021) or
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35 121 through unified model interoperable interfaces such as ESMF (Earth System
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38 122 Modeling Framework; Hill et al., 2004) and OpenMI (Open Modeling Interface;
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41 123 Harpham et al., 2019). However, the flexibility of such model-level integration may
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44 124 be limited by the inherent fixed model structures of each integrated model.
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47 125 Furthermore, different interoperable interfaces often need to be separately
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50 126 implemented for each model, posing considerable challenges for hydrologists lacking
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53 127 programming expertise.

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55 128 To further enhance the flexibility of spatially varying model structures, another
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58 129 idea involves internally constructing and integrating multiple model structures within
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1 130 a single modular hydrological modeling framework such as FLEX-Topo
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4 131 (Topography-driven Flux Exchange hydrological model; Gao et al., 2014; Savenije,
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6 132 2010), SUPERFLEX (SUPER Flux Exchange hydrological model; Fenicia et al.,
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9 133 2011, 2016), and airGR (Suite of GR Hydrological Models; Coron et al., 2017;
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12 134 Thébault et al., 2024). These frameworks allow the assignment of different model
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15 135 structures to different modeling areas within the watershed with only minimal
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18 136 changes to configuration or source code (Thébault et al., 2024). Nevertheless, these
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21 137 frameworks are fundamentally limited to conceptual representations of hydrological
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24 138 processes, with simulation units typically being subbasins and HRUs. Therefore, these
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27 139 state-of-the-art flexible frameworks primarily support lumped or semi-distributed
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30 140 conceptual model structures and thus face a significant challenge in integrating
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33 141 physically-based simulation algorithms and spatially explicit distributed simulation
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36 142 units.

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38 143 To address this limitation, we propose a new and practically significant spatially
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41 144 hybrid modeling approach. This approach enables the flexible combination of
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44 145 compatible spatial units and simulation algorithms to construct distinct model
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47 146 structures for individual subbasin within the watershed, such as lumped conceptual
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50 147 and fully-distributed semi-physically-based model structures. Subbasin-level model
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53 148 structures are coherently integrated through channel routing to perform watershed-
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56 149 scale simulation. The remainder of this paper is organized as follows: Section 2
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59 150 outlines the design of the proposed approach and its implementation within the
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1 151 SEIMS (Spatially Explicit Integrated Modeling System; Zhu et al., 2019) framework,
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4 152 Section 3 presents an experimental case study, followed by discussion in Section 4,
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6 153 and the conclusion in Section 5.
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9 154 **2 Method design and implementation**

10 155 **2.1 Basic idea**

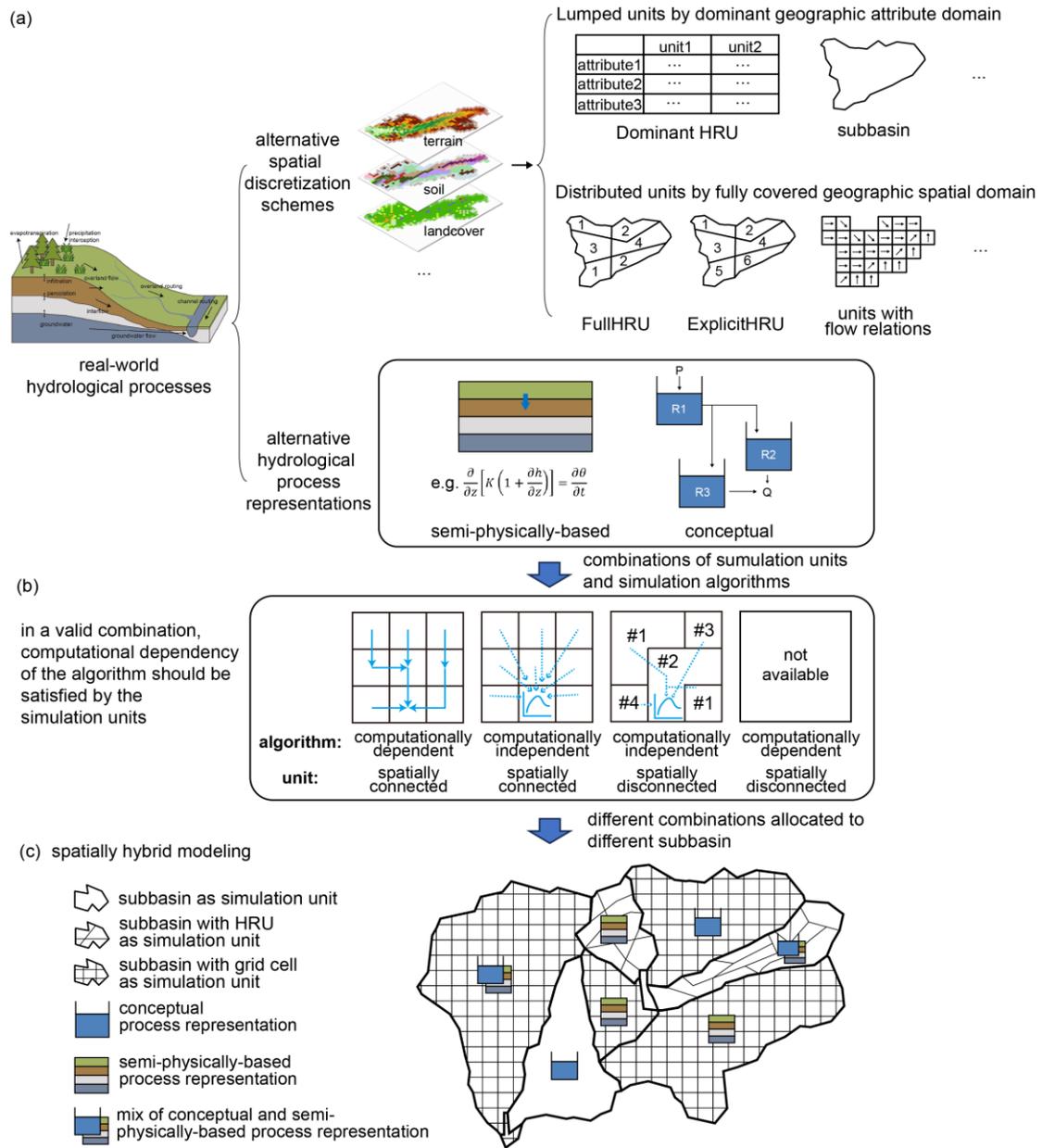
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16 156 The basic idea of the proposed spatially hybrid hydrological modeling approach
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19 157 is employing a divide-and-conquer strategy to construct and execute different model
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22 158 structures for relatively independent areas within the watershed. A watershed can be
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25 159 divided into distinct closed catchment areas connected by the channel network, i.e.,
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28 160 subbasins. Thus, subbasins are naturally suited to serve as the divide-and-conquer
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31 161 modeling units, each configurable with an individual model structure. Therefore, the
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34 162 core design of this approach involves two key aspects: 1) flexibility in constructing
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37 163 model structures, and 2) the capability to integrate these subbasin-specific model
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40 164 structures to perform watershed-scale simulation.
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42 165 The idea in designing a flexible model structure is to ensure compatibility
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45 166 between the spatial discretization scheme (i.e., simulation units) and hydrological
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48 167 process representations (i.e., simulation algorithms). Both can be classified according
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51 168 to their dependencies, which determine their compatibility (Figure 1a). For simulation
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54 169 units, dependency (or connectivity) refers to whether spatially explicit relationships
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57 170 exist between units. For simulation algorithms, computational dependency refers to
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60 171 whether a simulation on one unit relies on outputs (e.g., water flows or state variables)
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1 172 from adjacent or upstream units (Figure 1b). Therefore, a valid combination requires
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4 173 that computational dependencies of selected simulation algorithms are satisfied by the
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6 174 spatial connectivity of the simulation units (Figure 1b). For instance, a conceptual
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9 175 reservoir-based surface runoff generation algorithm combined with a physically-based
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12 176 kinematic wave-based surface routing can be applied to grid cells, where
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15 177 computational dependency is satisfied by the flow directions between grid cells.
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18 178 Another example is the combination of a physically-based Green-Ampt surface runoff
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21 179 generation algorithm with a conceptual unit hydrograph-based surface routing
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24 180 algorithm for application on HRUs (Craig, 2020), which requires no computational
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27 181 dependency. It is crucial to note that while computational dependency is a
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30 182 fundamental constraint, the ultimate rationality of a customized model structure relies
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33 183 on the modeler's hydrological knowledge and application-specific requirements.

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35 184 Each subbasin model simulates hillslope processes and the channel routing
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38 185 process. In this study, it is assumed that mass and energy are exchanged between
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41 186 subbasins only through the channel network. Thus, the channel routing components
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44 187 within each subbasin model structure (e.g., the Muskingum method) are used to
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47 188 couple with their upstream and downstream subbasins, integrating all subbasin model
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50 189 structures into a comprehensive watershed model (Figure 1c). The channel routing of
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53 190 a downstream subbasin relies on outputs from its upstream channels. This dependency
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56 191 implies that the hillslope processes of each subbasins can be parallelized to improve
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59 192 computing efficiency, while channel routing processes of each subbasins must be
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193 executed sequentially according to the upstream–downstream network structure (Liu
 194 et al., 2016).



195
 196 Figure 1. Design of the proposed spatially hybrid hydrological modeling approach. a)
 197 defining model structure by spatial discretization schemes and hydrological process
 198 representations; b) supporting compatible combinations of simulation units and
 199 simulation algorithms; and c) enabling the spatial varying allocation and integration
 200 of different model structures at the subbasin level.

2.2 Design of the spatially hybrid hydrological modeling approach using the SEIMS framework

Following the above basic idea, the proposed spatially hybrid hydrological modeling approach was designed using the SEIMS framework for its flexible modular design and two-level parallelization strategy (Liu et al., 2016; Zhu et al., 2019). To support the flexible construction of model structures, the spatial discretization schemes of SEIMS were expanded (Figure 2a; Section 2.2.2), the simulation modules of hillslope processes were upgraded to accommodate compatible simulation units (Figure 2b; Section 2.2.2), and the execution of SEIMS-based model was upgraded to allocate different model structures to subbasins and executed in an integrated way (Figure 2c; Section 2.2.3).

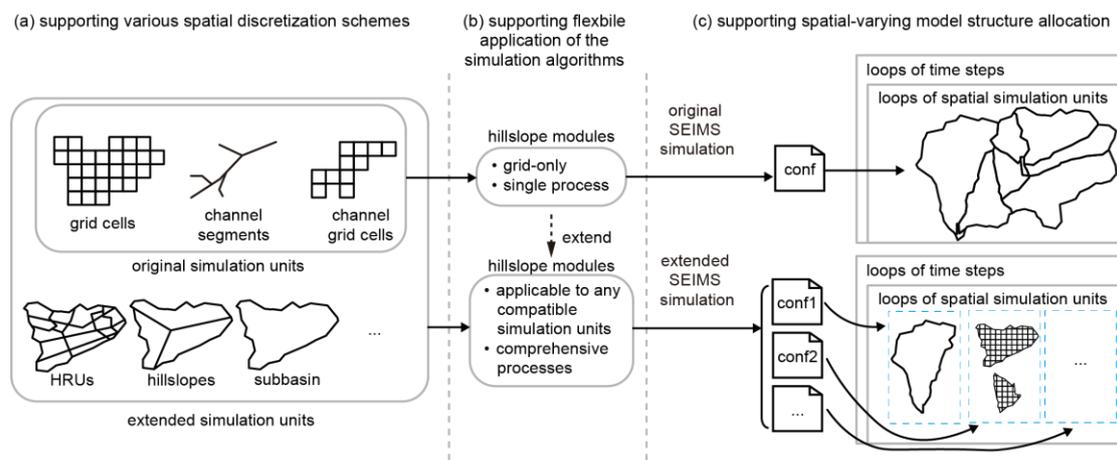


Figure 2. Overall design of extending the SEIMS (Spatially Explicit Integrated Modeling System) to implement the proposed spatially hybrid hydrological modeling approach: a) supporting various spatial discretization schemes; b) supporting applying simulation modules of hillslope processes to compatible simulation units; and c) supporting spatial varying model structure allocation on subbasins to construct a watershed model.

1 220 **2.2.1 Brief introduction to SEIMS**

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4 221 As a hydrological modeling framework, SEIMS provides users with the
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7 222 complete toolchain to preprocess modeling data, construct and execute the SEIMS-
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10 223 based model, analyze the parameter sensitivity, and so on (Liu et al., 2021; Zhu et al.,
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13 224 2019). More importantly, SEIMS supports developers in adding simulation modules
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16 225 of hydrologic processes following three developing principles: the spatial
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19 226 discretization scheme, the modular modeling design, and the parallelization strategy.

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21 227 **(1) Spatial discretization scheme**

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24 228 SEIMS primarily adopts a two-level spatial discretization scheme, i.e., the
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27 229 “subbasin–basic simulation unit.” In the stage of preprocessing modeling data, SEIMS
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30 230 delineates the watershed into subbasins, each with one channel. Within each subbasin,
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33 231 the basic simulation units (i.e., grid cells in the original version) are delineated and
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36 232 organized as a one-dimension array. Most simulation modules of hillslope processes
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39 233 use grid cells as simulation units, while a few modules use subbasin units such the
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42 234 linear reservoir algorithm for the groundwater process. Modules of channel routing
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45 235 processes use channel lines or grid cells as simulation units. All parameters required
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48 236 by the modules on these simulation units are prepared using the data preprocessing
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51 237 tools. The grid cells are layered based on the flow direction of each cell (i.e., the
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54 238 single flow direction in the original version; Liu et al., 2014). The layers are used in
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57 239 modules of surface or subsurface flow routing processes to determine the simulation
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60 240 sequences of grid cells. The same idea is used for layering subbasins and channels.

1 241 **(2) Modular modeling design**

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3 242 A SEIMS-based hydrological model is constructed by the SEIMS main program
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6 243 and several user-configured modules of hydrologic processes in a text-based format
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9 244 rather than hard-coded. The main program controls the time loop of the simulation
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12 245 and repeatedly invokes the selected modules in sequence. Each module is responsible
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15 246 for the simulation of one or several hydrologic processes on corresponding simulation
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18 247 units in different simulation orders. For instance, the potential evapotranspiration
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21 248 module executes on every grid cell in an arbitrary order, the subsurface flow (also
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24 249 referred to as interflow) routing module executes on every grid cell by layering orders
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27 250 (Liu et al., 2014), and the channel routing module executes on each channel by
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30 251 upstream–downstream orders.

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32 252 In SEIMS, each module should handle its required simulation conditions. Users
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35 253 should decide the integration feasibility of the module combinations to meet the
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38 254 requirements of each module, such as the boundary conditions when calculating
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41 255 vertical and lateral water movements. For instance, to implement a semi-physically-
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44 256 based interflow module based on the Darcy’s Law, the boundary condition could be
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47 257 set either with the hydraulic gradient equal to the slope at each cell, or dynamically
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50 258 determined using the outputs of adjacent cells provided by other modules.

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52 259 Each module inherits from a standard module interface, including the definitions
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55 260 of input data, parameters, and output data, and the core execution code. The input data
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58 261 and parameters of one module can be either read from the database created during the

1 262 preprocessing stage or referred from other modules during the runtime. Each module
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4 263 itself initializes its output data. In this way, during the execution of the SEIMS-based
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6 264 model, the same variable is shared among modules.

9 265 **(3) Two-level parallelization strategy**

11 266 The first level of the two-level parallelization strategy dispatches the simulation
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15 267 of subbasins to different computing processes (or nodes) through the Message Passing
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18 268 Interface (MPI) (Liu et al., 2016). The second level is achieved within each subbasin
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21 269 by dispatching the simulations of grid cells without mutual dependencies to
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24 270 computing threads via Open Multi-Processing (OpenMP) (Liu et al., 2014). Based on
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27 271 the design of the modular structure, variables required to be communicated among
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30 272 subbasins only need to be defined in the metadata of each module and will be handled
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33 273 by the SEIMS main program (Zhu et al., 2019). The two-level parallelization strategy
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36 274 provides the potential to implement the proposed approach in this study that applies
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39 275 different model structures to distinct subbasins.

41 276 **2.2.2 Constructing model structures with diverse simulation units and algorithms**

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45 277 To enable the flexible construction of model structures based on SEIMS, the
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48 278 diversity of both simulation units and simulation algorithms should be guaranteed. For
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51 279 simulation units, hydrological response units (HRUs) were added as simulation units
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54 280 of hillslope processes in SEIMS, including DominantHRU (defined by dominant
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57 281 combination of soil types, land use types, and slope classes), FullHRU (one HRU may
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60 282 consist of several spatially separated areas), and ExplicitHRU (one HRU is a spatially

1 283 independent and contiguous area) (Figure 1a; Figure 2a). The physically-based
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4 284 parameters can be derived from actual properties of the soil, landuse/landcover, and
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6 285 other spatial data, while conceptual parameters are directly specified by lookup tables
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9 286 of soil and land-use. For instance, the storage capacity of an HRU can be represented
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12 287 by the soil depth associated with that specific HRU area, while some experimental
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15 288 coefficients could be specified in the lookup tables. Under such an HRU discretization
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18 289 scheme, the hillslope unit and subbasin unit can be regarded as specific
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21 290 DominantHRUs. Therefore, the extended SEIMS will offer the flexibility of utilizing
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24 291 subbasins, hillslopes, HRUs, and grid cells as simulation units of hillslope processes.

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26 292 For hydrological process representations, SEIMS was initially designed to
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29 293 primarily implement the simulation algorithm of one hydrological process as a single
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32 294 module, while it lacked a conceptual model capable of considering multiple
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35 295 hydrological processes comprehensively (Figure 2b). In this study, lumped conceptual
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38 296 models were designed to be integrated as two separate hillslope modules: (1) a surface
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41 297 runoff module that generate runoff at each basic simulation unit; and (2) a surface
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44 298 routing module that convey runoff to the subbasin outlet. For instance, the
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47 299 representative widely used conceptual model, GR4J (Perrin et al., 2003), is suitable to
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50 300 be incorporated as a surface runoff module, coupled with a separate routing module
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53 301 such as a simple accumulation method. Such a module could be applied to any
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56 302 simulation unit types stated above.

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58 303 To ensure the combination compatibility of simulation algorithms with
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1 304 simulation units applied to subbasins, the SEIMS module interface was extended to
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4 305 mark its computational dependency requirement. A computationally dependent
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6 306 simulation module is exclusively applicable to grid cells, and a computationally
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9 307 independent module is applicable to any types of simulation unit. A model structure
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12 308 containing any computationally dependent module can only be applied to the subbasin
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15 309 using grid cells as the basic simulation units.

18 310 **2.2.3 Allocating individual model structures to subbasins and integrating as a** 19 20 21 311 **spatially hybrid watershed model structure**

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24 312 An essential part of the spatially hybrid hydrological modeling approach is to
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27 313 enable the separate allocation and execution of individual model structures to different
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30 314 subbasins (referred to as subbasin models) within the watershed. All subbasin models
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33 315 are subsequently integrated as a watershed model. The requirements of flexible model
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36 316 configuration and subbasin-separate simulation are compatible with the text-based
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39 317 module configuration method and two-level parallelization strategy of SEIMS, but
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42 318 still need improvement. The module configuration file of SEIMS was extended to
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45 319 designate a model structure to the specific subbasin, and the adopted type of
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48 320 simulation units (Figure 3). In this way, the SEIMS main program was extended to
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51 321 read the configuration file for each subbasin dynamically, load simulation modules,
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54 322 and retrieve the modeling parameters according to the specified basic simulation
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57 323 units.

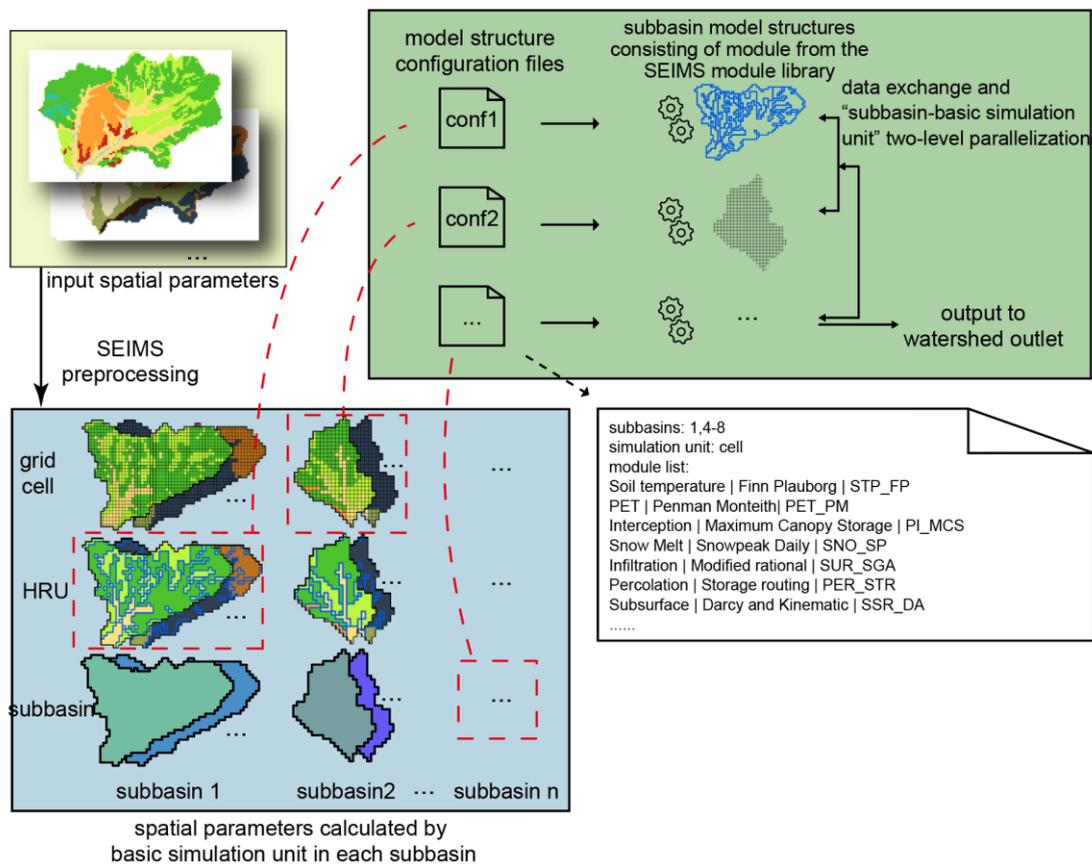
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59 324 The two-level parallelization strategy was improved in two aspects. The first is

1 325 task scheduling of the subbasin-level parallelization for load balancing handled by the
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4 326 SEIMS main program. The domain decomposition of subbasins was determined by
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6 327 upstream–downstream relationships between subbasins and the number of basic
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9 328 simulation units of each subbasin in the runtime, rather than using the numbers of grid
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12 329 cells of each subbasin in the preprocessing stage in the original version of SEIMS
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15 330 (Liu et al., 2016). The second aspect concerns the parallelization at the basic
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18 331 simulation unit level, which is handled in computationally independent modules. This
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21 332 is inherently supported and easy to implement, as the OpenMP for-loop can also apply
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24 333 to newly added basic simulation units such as HRUs.
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30 335 **2.3 Implementation**

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33 336 Based on the above method design, the implementation of the proposed approach
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36 337 with SEIMS involves modifications to the data preprocessing tools, SEIMS main
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39 338 programs, and SEIMS module library. The data preprocessing tools mainly include a
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42 339 collection of Python scripts and C++ programs to delineate spatial units at different
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45 340 scales (e.g., subbasins and hillslopes), extract spatial parameters of spatial units, and
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48 341 create watershed modeling database (Zhu et al., 2019). In this study, a configurable
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51 342 tool was implemented to support the delineation and parameterization of HRU-based
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54 343 spatial units. This tool allows users to specify environmental variables (e.g., the
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57 344 default land-use and soil types) to delineate HRU-based units. This tool also allows
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60 345 specifying base spatial units and delineation strategy to generate different types of
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346 HRU such as taking subbasins or hillslopes as base units to generate DominantHRU
 347 or FullHRU. Therefore, the data preprocessing tools could prepare various types of
 348 spatial unit and associated spatial parameters for each subbasin according to the user-
 349 defined configuration file to meet various hydrological modeling scenarios. The
 350 procedures of spatially hybrid modeling with SEIMS are depicted in Figure 3.



352 Figure 3. Procedures of spatially hybrid modeling with the SEIMS (Spatially Explicit
 353 Integrated Modeling System)

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 355 The SEIMS main program, written in C++, was extended to read the
 356 configuration file for each subbasin, load simulation modules and retrieve the
 357 modeling parameters according to the specified simulation units. Simulation modules

1 358 of the SEIMS module library were updated to declare the type of computational
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4 359 dependency. New conceptual simulation modules were added such as GR4J for
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6 360 simulating hillslope runoff generation. With the above briefly introduced
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9 361 implementation of the proposed approach, SEIMS can now support users in
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12 362 constructing spatially hybrid model structures for considering spatial heterogeneity of
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15 363 the watershed, where each subbasin may be simulated with different combinations of
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18 364 conceptual or semi-physically-based simulation modules and spatial units.
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21 365 Consequently, more than 40 modules are available after the implementation of this
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24 366 study (detailed in Table S1 of the supplementary material), supporting simulation of
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27 367 processes including snowfall, atmospheric deposition, snow balance, interception, soil
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30 368 temperature, glacier, surface runoff, evapotranspiration, infiltration, percolation,
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33 369 interflow, groundwater, channel routing, plant growth, and soil erosion.

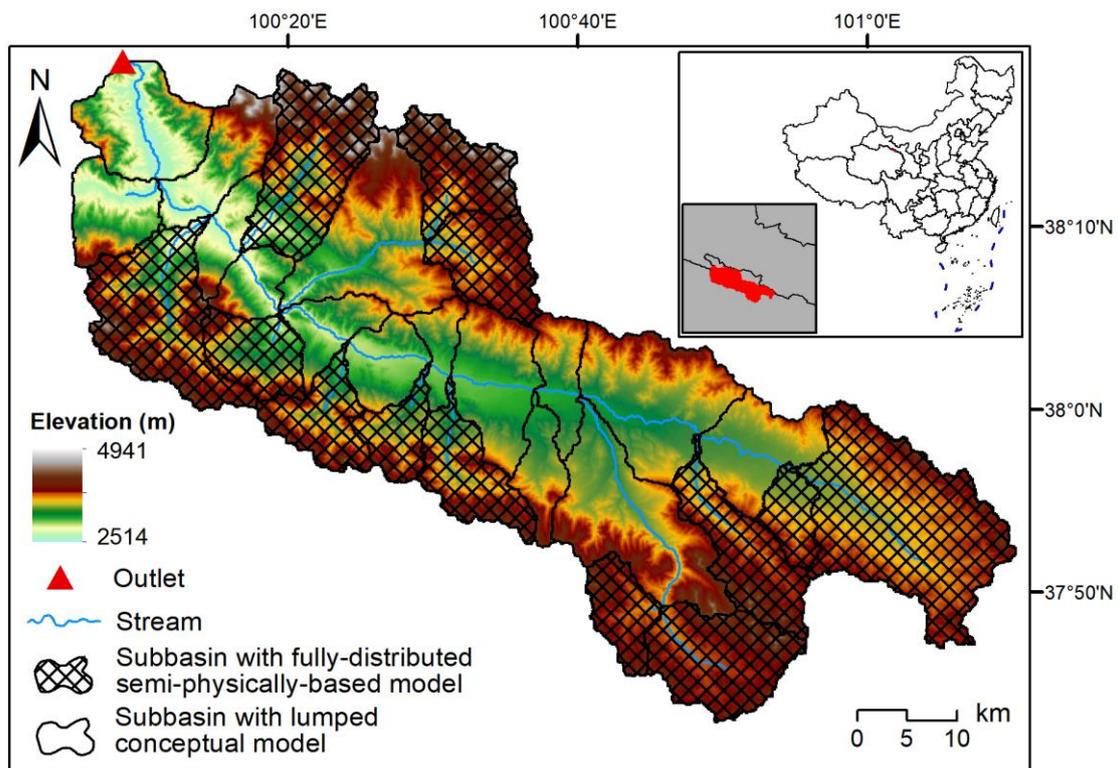
34
35 370 The SEIMS is open-source on GitHub (<https://github.com/lreis2415/SEIMS>) and
36
37
38 371 is under continuous development.

372 **3 Case study**

373 **3.1 Study area and data**

374 The case study area is the Babao River watershed at Qilian, Qinghai Province,
375 China (Figure 4), a high-altitude, cold and mountainous region with an area of
376 approximately 2,511 km². The average elevation is 3,565 m, and the region features
377 glacier, snow cover and frozen soil. The data used in the case study are listed below.
378 The MERIT DEM with the resolution of 90 m (Yamazaki et al., 2017) was selected

1 379 for watershed delineation and calculation of terrain attributes. Meteorological data
2
3
4 380 was obtained from the China Meteorological Assimilation Driving Datasets for the
5
6 381 SWAT model (CMADS) version 1.2, provided in the form of approximately 0.125°
7
8
9 382 resolution gridded station data (Meng et al., 2017). Land cover data was derived from
10
11
12 383 GLOBELAND30 with a resolution of 30 m (<http://www.globallandcover.com>). Soil
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15 384 attribute data was sourced from the Harmonized World Soil Database (HWSD)
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17
18 385 Version 2.0 (Nachtergaele, 2023). The observed daily discharge of the outlet gauge
19
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21 386 from 2013 to 2018 was obtained from the National Hydrological Yearbook.



387
388 Figure 4. The Babao River watershed and the spatial constitution of the spatially
389 hybrid watershed model.

3.2 Experimental design

To verify the integration feasibility and simulation effectiveness of the proposed spatially hybrid hydrological modeling approach and its implementation using SEIMS, we constructed one fully-distributed semi-physically-based model structure (section 3.2.1) and one lumped conceptual model structure (section 3.2.2) to build one spatially hybrid model and two spatially consistent models (section 3.2.3) for comparison after parameter calibration (section 3.2.4). It is worth noting that the watershed models constructed in this comparative experiment are not intended to model the hydrological processes in the study precisely, nor to enumerate the possible model structures applicable as illustrated in Figure 1.

3.2.1 The fully-distributed semi-physically-based model structure

The fully-distributed semi-physically-based model structure uses grid cells as the basic simulation units and encompasses simulation algorithms based on physical laws. For instance, the algorithm for percolation is calculated as the product of hydraulic conductivity and the gradient of the hydraulic potential, and the interflow is calculated from the kinematic approximation of Darcy's Law and, with the hydraulic gradient equal to the slope at each grid cell. These two simulation algorithms are the same as those used in WetSpa (Water and Energy Transfer between Soil, Plants and Atmosphere; Liu and Smedt, 2004). Table 1 lists the considered hydrological processes and their simulation algorithms associated with key parameters for calibration. The physically based simulation algorithms primarily utilize spatial

411 parameters with clear physical meaning, such as soil hydraulic conductivity. Besides,
 412 this model structure also includes conceptual simulation algorithms such as the
 413 surface routing module based on unit hydrograph (Table 1).

414

415 Table 1. Simulation algorithms adopted for the fully-distributed semi-physically-
 416 based model structure and the parameters involved in model calibration

Hydrological process	Simulation algorithm	Parameters involved in calibration
Potential evapotranspiration	Penman-Monteith	K_pet (correction factor)
Interception	Maximum canopy storage	Interc_max (maximum interception storage), Interc_min (minimum interception storage), Pi_b (interception storage capacity exponent)
Glacier	HBV method	-
Snow melt	HBV method	T_rain_snow_delta (rain/snow mixture temperature range)
Infiltration and surface runoff	Modified coefficient method	K_run (runoff exponent), P_max (maximum precipitation corresponding to potential runoff coefficient), Runoff_co (potential runoff coefficient)
Depression	Linsley	Depression (depression storage capacity)
Percolation	Storage routing	-
Interflow	One-dimension kinematic wave	Ki (interflow scale factor)
Actual evaporation	Thornthwaite and Mather method	-
Plant growth	Simplified EPIC	-
Groundwater	Linear reservoir	Base_ex (baseflow recession exponent), df_coef (deep percolation coefficient), gwmax (maximum ground water storage), Kg (baseflow recession coefficient)
Surface routing	Geomorphology-based unit hydrograph	-
Channel routing	Muskingum	MSK_X, MSK_K

1 417 **3.2.2 The lumped conceptual model structure**

2
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4 418 The lumped conceptual model structure adopts DominantHRUs as the simulation
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6
7 419 units and GR4J as the main simulation algorithm (Table 2). The Hargreaves method is
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10 420 used to estimate the potential evapotranspiration for the GR4J. The GR4J receives the
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13 421 potential evapotranspiration and precipitation as input to generate the hillslope runoff
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16 422 at the outlet of the simulation unit (i.e., DominantHRUs), which implicitly generalizes
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18
19 423 the internal surface routing process. The runoff from all DominantHRUs is then
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21
22 424 aggregated to the subbasin outlet using a simple accumulation method. The inter-
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24
25 425 subbasin channel routing process is modeled using the Muskingum algorithm. In this
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27
28 426 case study, each subbasin corresponds to a single DominantHRU, defined by the land
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30
31 427 cover and soil type combination that occupies the largest area within the subbasin. As
32
33 428 a result, GR4J can also be applied without an additional surface routing module.

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35 429

36
37 430 Table 2. Simulation algorithms adopted for the lumped conceptual model structure
38
39 431 and the parameters involved in model calibration

Hydrological process	Simulation algorithm	Parameters involved in calibration
Potential evapotranspiration	Hargreaves	-
Surface runoff	GR4J	SOILTHICK, X2, X3, X4
Surface routing	Simple accumulation method	-
Channel routing	Muskingum	MSK_X, MSK_K

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1 433 **3.2.3 One spatially hybrid and two spatially consistent model structures for**
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4 434 **comparison**

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7 435 Ideally, physically based models can better simulate hydrological processes with
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10 436 explicit physical representations. However, due to practical limitations, such as
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13 437 incomplete understanding of hydrological processes and the difficulty of
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16 438 implementing complex simulation algorithms and deriving essential modeling data, a
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19 439 conceptual model can serve as a valuable complement. In this case study, we
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22 440 constructed a spatially hybrid model structure (referred to as the HybM; Figure 4) by
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25 441 combining both fully-distributed semi-physically-based (see Section 3.2.1) and
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28 442 lumped conceptual models (see Section 3.2.2). We considered the fully-distributed
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31 443 semi-physically-based model structure more suitable for mountainous subbasins (i.e.,
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34 444 13 subbasins shown in Figure 4), while applying the lumped conceptual model
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37 445 structure to gently sloping subbasins (i.e., 16 subbasins shown in Figure 4).

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39 446 For comparison with the HybM, we also constructed two spatially consistent
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42 447 model structures: PhyM, which uses only the fully-distributed semi-physically-based
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45 448 model structure, and ConM, which uses only the lumped conceptual model structure.

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48 449 **3.2.4 Comparative experiments**

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51 450 The comparative experiments primarily aim to assess the post-calibration
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54 451 performance of the three watershed model structures. Parameters for model
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57 452 calibration were chosen through prior sensitivity analyses (listed in Table 1 and Table
58
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60 453 2). In principle, the proposed approach adheres to the atomistic forward modeling

1 454 concept: each subbasin-level model structure should be determined based on the
 2
 3 455 appropriate representation of hydrological processes and available data, ensuring a
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 6 456 physically reasonable and practically executable watershed model. Ideally, model
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 9 457 calibration should utilize multi-scale observations (e.g., at subbasin and watershed
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 12 458 outlets) to avoid relying solely on the inverse calibration assumption that successfully
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 15 459 watershed-scale reproduction validates subbasin-scale model structures.

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 18 460 In the absence of multi-scale observations, it is a common practice in current
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 21 461 hydrological modeling to calibrate based solely on the watershed outlet data. In this
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 24 462 case study, only the watershed outlet discharge data was available. To partially
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 27 463 mitigate this limitation, we designed two calibration strategies (i.e., universal and
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 30 464 regional). The universal calibration strategy involves a uniform adjustment of
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 33 465 parameters for all subbasin models, while the regional calibration strategy applies
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 36 466 separate parameter sets to models of mountainous subbasins and gently sloping
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 39 467 subbasins. Consequently, the HybM was calibrated using the regional strategy, while
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 41
 42 468 both PhyM and ConM were calibrated using both universal and regional strategies
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 44 469 (Table 3).

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 48
 49 471 Table 3. Experiments of parameter calibration with different calibration strategies for
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 51 472 different watershed model structures

Calibration strategy	Name of experimental cases		
	ConM	PhyM	HybM
Universal	ConM1	PhyM1	-
Regional	ConM2	PhyM2	HybM

473 All three models were executed over a six-year simulation (January 1, 2013–
 474 December 30, 2018). The first year (2013) served as a warm-up period, followed by a
 475 three-year calibration period (2014–2016) and a two-year validation period (2017–
 476 2018). Model performance was evaluated using three standard indices: the Nash-
 477 Sutcliffe coefficient (NSE; Eq. 1), root mean square error-standard deviation ratio
 478 (RSR; Eq. 2), and percent bias (PBIAS; Eq. 3).

$$NSE = 1 - \frac{\sum_{i=1}^n (O_i - P_i)^2}{\sum_{i=1}^n (O_i - \bar{O})^2} \quad (1)$$

$$RSR = \frac{\sqrt{\sum_{i=1}^n (O_i - P_i)^2}}{\sqrt{\sum_{i=1}^n (O_i - \bar{O})^2}} \quad (2)$$

$$PBIAS = \frac{\sum_{i=1}^n (O_i - P_i) \times 100}{\sum_{i=1}^n O_i} \quad (3)$$

479 where O_i and P_i are i -th observed value and predicted value, respectively. \bar{O} is the
 480 averaged observed value, and n is the size of simulated time series.

481 The parameter calibration experiments were conducted using the NSGA-II (non-
 482 dominated sorting genetic algorithm-II; Deb et al., 2002) integrated in the SEIMS
 483 framework (Zhu et al., 2019). The optimization parameters setting for each calibration
 484 experiment included a generation of 25 and a population of 360. The multi-objective
 485 optimization aimed to maximize NSE and minimize RSR and absolute value of
 486 PBIAS. Experiments were carried out on a server equipped with a 2.70 GHz Intel
 487 Xeon Gold 6150 dual CPU (36 cores).

3.3 Experimental results and discussion

The integration feasibility of the constructed spatially hybrid watershed model structure is demonstrated by the successful results of the comparative experiments. To evaluate the simulation effectiveness of the proposed approach, one set of calibrated parameters of each experiment (detailed in Table S2 of the supplementary material) was selected for comparison, analyzing both model performance metrics and the hydrographs.

3.3.1 Model performance metrics of calibrated model structures

Table 4 lists the model performance metrics for both calibration and validation periods of calibrated model structures under the universal and regional calibration strategies.

Table 4. Model performance metrics of calibrated model structures under different calibration strategies (PhyM1 and PhyM2 represent fully-distributed semi-physically-based model structure using universal and regional calibration, respectively; ConM1 and ConM2 represent lumped conceptual model structure using universal and regional calibration, respectively; HybM represents the spatially hybrid model structure using regional calibration; NSE: Nash-Sutcliffe coefficient; RSR: root mean square error-standard deviation ratio; PBIAS: percent bias; values are reported as calibration/validation)

Experimental case	Performance metrics of calibration/validation periods		
	NSE	RSR	PBIAS (%)
PhyM1	0.58/0.60	0.64/0.63	10.63/2.58
PhyM2	0.59/0.60	0.64/0.63	17.34/13.90
ConM1	0.41/0.32	0.77/0.82	2.89/-42.87
ConM2	0.45/0.50	0.74/0.71	10.96/-21.03
HybM	0.72/0.60	0.53/0.63	10.02/-18.28

1 507 The spatially consistent fully-distributed semi-physically-based models (PhyM)
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3
4 508 exhibited similar NSE (~0.59) and RSR (~0.63) in both calibration and validation
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6 509 periods, regardless of the calibration strategy. However, PhyM1 (universal calibration)
7
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9 510 yielded a lower PBIAS than PhyM2 (regional calibration). While regional calibration
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12 511 generally allows for more flexible parameter adjustments across heterogeneous
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15 512 subbasins, the higher overall bias observed in PhyM2 may be attributed to the lack of
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18 513 internal subbasin observation data. Without such internal constraints, regional
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21 514 calibration may lead to spatially inconsistent simulation errors that accumulate toward
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24 515 the watershed outlet, thereby increasing the overall bias. In contrast, universal
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27 516 calibration often benefits from compensatory effects among subbasins, resulting in a
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30 517 lower overall bias.

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32 518 Unlike the effect of calibration strategies on PhyM, the spatially consistent
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35 519 lumped conceptual models (ConM) showed moderate improvements under the
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38 520 regional calibration strategy (ConM2) compared to the universal calibration strategy
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41 521 (ConM1). Specifically, for calibration/validation, the NSE increased from 0.41/0.32 to
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44 522 0.45/0.50 and the RSR decreased from 0.77/0.82 to 0.74/0.71. For PBIAS, ConM2
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47 523 yielded 10.96%/-21.03%, showing improved validation bias compared to ConM1
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50 524 (2.89%/-42.87%), despite a higher calibration-period bias. Conceptual models like
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53 525 ConM, with simple model structure and few parameters, are often considered easier to
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55
56 526 calibrate and potentially perform better under regional calibration. However, in this
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59 527 study, both ConM1 and ConM2 exhibited relatively low NSE and high RSR values in
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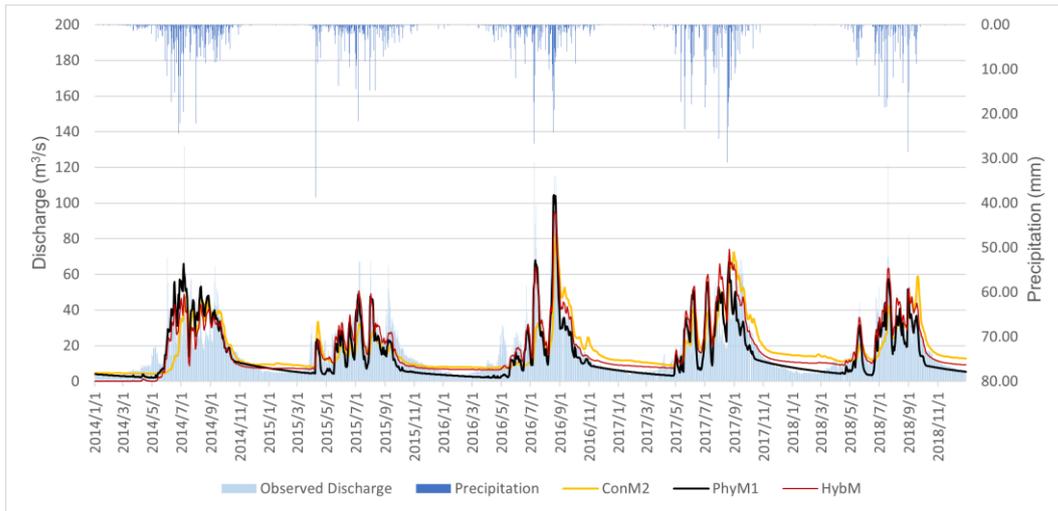
1 528 both periods, indicating the limited capability of the GR4J-based conceptual model
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3 529 structure to represent the hydrological processes in this mountainous watershed,
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6 530 particularly because it does not account for glacier and snowmelt processes. The
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9 531 noticeable decline in validation performance for both ConM1 and ConM2 (e.g., the
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12 532 substantial shift in PBIAS from overestimation to underestimation) further reflects the
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15 533 limited generalization of ConM. The contrasting results between ConM and PhyM
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18 534 may be attributed to their structural differences in handling spatial variability. ConM
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21 535 relies more on regional calibration to consider spatial heterogenous across subbasins
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23
24 536 by adjusting its less physically constrained parameters. In contrast, PhyM captures
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27 537 spatial heterogeneity directly through spatially distributed input data and parameters,
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30 538 implying a low dependency on calibration strategies to represent spatial variations in
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33 539 hydrological processes.

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35 540 The spatially hybrid model structure (HybM) demonstrated a balanced model
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38 541 performance by leveraging the strengths of both PhyM and ConM, achieving the
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41 542 highest NSE (0.72/0.60) and lowest RSR (0.53/0.63) among all experimental cases.
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44 543 However, its PBIAS (10.02%/-18.28%) was lower than (i.e., better overall bias) that
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47 544 of both ConM cases but higher than (i.e., worse overall bias) both PhyM cases. This
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50 545 reflects that the spatially hybrid model structure provides a promising compromise
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53 546 between modeling flexibility and simulation accuracy. That is, while HybM combined
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56 547 the calibration responsiveness of ConM and the process realism of PhyM, it also
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59 548 inherited some of their respective limitations, such as bias associated with the
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1 549 conceptual representation. Overall, these findings highlight the importance of careful
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4 550 model structure construction and the necessity of multi-site and long-term observation
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6 551 data to effectively constrain calibration and fully realize the benefits of the proposed
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9 552 approach.

10 11 12 553 **3.3.2 Hydrographs of calibrated model structures**

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16 554 Figure 5 compares the outlet hydrographs simulated by PhyM1, ConM2, and
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18 555 HybM against the observed discharge. During the dry seasons (typically November to
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21 556 April), when baseflow dominates streamflow, PhyM1 better simulates baseflow
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23 557 magnitudes, benefiting from its process-based representation of hydrological
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26 558 processes. In contrast, ConM2 tends to overestimate baseflow, likely due to its
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28
29 559 simplified recession formulation and lack of explicit representation of groundwater
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31
32 560 storage and release dynamics. During the wet season, characterized by high-intensity
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35 561 precipitation events, PhyM1 produces sharper rising and falling limbs, generally
36
37
38 562 capturing the dynamics of observed peaks and recessions. This behavior reflects the
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40
41 563 physical realism of PhyM in representing rapid surface and subsurface responses in
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43
44 564 the study area. Conversely, ConM2 often either overestimates or underestimates peak
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47 565 flows, likely due to limitations in its conceptual routing and storage structure.
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567 Figure 5. Simulated outlet hydrographs of best-performing cases for each model
 568 structure: PhyM1 (fully-distributed semi-physically-based model structure using
 569 universal calibration strategy), ConM2 (lumped conceptual model structure using
 570 regional calibration strategy), and HybM (spatially hybrid model structure using
 571 regional calibration strategy)

572

573 The spatially hybrid model (HybM) integrates the strengths of both model
 574 structures. HybM more accurately captures peak magnitudes and recession processes
 575 throughout the simulation period, exhibiting rising limbs similar to PhyM1 and a
 576 moderated baseflow response that is less overestimated than in ConM2. Although
 577 HybM still overestimates baseflow compared to observations, its improved flood
 578 response and recession simulation suggest a more balanced simulation across different
 579 hydrological conditions.

580 4 Discussion

581 The case study results have demonstrated the integration feasibility and
 582 simulation effectiveness of the proposed spatially hybrid hydrological modeling

1 583 approach. The results reflect an inherent trade-off in this approach, that is, enhanced
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4 584 flexibility and local adaptability often come at the cost of increased structural
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6 585 complexity and potential uncertainty. Nevertheless, the proposed approach offers a
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9 586 practical and adaptable framework to address spatial heterogeneity in diverse
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12 587 hydrological modeling application contexts, particularly where data availability and
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15 588 watershed characteristics vary across space. This section further discusses its broader
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18 589 implications, focusing on the significance of the advantages of SEIMS in
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21 590 implementing the proposed approach and limitations and future directions.

22 23 24 591 **4.1 Advantages of SEIMS in implementing the proposed** 25 26 27 592 **approach**

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30 593 SEIMS effectively supports the implementation of the spatially hybrid
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33 594 modeling approach through its modular and subbasin-independent architecture,
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36 595 offering both modeling flexibility and computational efficiency.

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39 596 SEIMS's design enables model developers to reuse or adapt existing model
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42 597 codes as SEIMS modules to simulate individual or multiple hydrological processes.
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45 598 For instance, the GR4J module was adapted from the Raven framework (Craig, 2020).
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48 599 Newly developed modules can be flexibly combined with existing ones to construct
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51 600 various model structures. This capability is particularly valuable for large and
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54 601 heterogeneous watersheds where modeling application contexts vary across regions.
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57 602 For instance, in subbasins with limited spatial data, lumped conceptual components
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60 603 can be employed in place of data-intensive distributed model structures to ensure

1 604 model feasibility. It is worth noting that while SEIMS provides this flexibility, module
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4 605 developers are responsible for specifying input requirements and boundary conditions
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6 606 for their modules, and users should assess the hydrological rationality and integration
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9 607 feasibility of their combined model structure.

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12 608 Regarding computational efficiency, in the case study, the fully-distributed
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15 609 semi-physically-based model structure required approximately 75 hours to complete
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18 610 9000 model runs for the calibration experiment, whereas the lumped conceptual
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21 611 model structure took only about 1 hour. The spatially hybrid model, incorporating
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24 612 both structures, achieved a balanced efficiency of around 30 hours. This
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27 613 computational efficiency, enabled by the parallelization strategies of SEIMS, supports
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30 614 more scalable applications in large watersheds, particularly those involving extensive
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33 615 model runs such as uncertainty analysis, parameter calibration, and scenario analysis.

34 35 36 616 **4.2 Limitations and future directions**

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39 617 This section highlights key limitations and outlines future directions to enhance
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42 618 its practical applicability. First, the successful application of the proposed approach
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45 619 implemented within SEIMS largely depends on the breadth and quality of its module
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48 620 library. Though extensible, the module library lacks modules for several critical
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51 621 hydrological processes (e.g., frozen soil dynamics) and specific representations (e.g.,
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54 622 fully-physically-based simulation algorithms that tightly couple hillslope hydrological
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56
57 623 processes). Recent research efforts continue to expand SEIMS module library to
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60 624 support more complex hydrological processes, for instance, by modeling

1 625 groundwater–surface water interactions between alpine runoff and alluvial aquifers in
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4 626 high mountain watersheds (Liu et al., 2025). We plan to gradually implement a variety
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6 627 of fully-physically-based simulation modules, such as a 2D hillslope overland flow
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9 628 module inspired by FullSWOF (Full Shallow-Water equations for Overland Flow;
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12 629 Delestre et al., 2017), and a fully coupled 3D module of overland flow, subsurface
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15 630 flow, and groundwater flow following approaches from SHUD (Shu et al., 2020) and
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18 631 ParFlow (Kollet and Maxwell, 2006). We encourage model developers to contribute
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21 632 to the enrichment of the SEIMS module library, either by transplanting existing code
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24 633 or developing new modules based on established hydrological equations and the
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27 634 SEIMS module template.

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29 635 Second, while the subbasin-independent architecture of SEIMS is a key feature
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32 636 to implement the proposed approach, it also raises concerns about inter-subbasin
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35 637 connectivity for more complex modeling circumstances. For instance, when multiple
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38 638 flow direction algorithms are used to construct the hillslope flow routing network,
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41 639 flow paths may cross subbasin boundaries and the connectivity between subbasins
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44 640 becomes more frequent than only connected by the channel network. Future work
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47 641 should explore flexible inter-subbasin connectivity mechanisms to consider more
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50 642 realistic hydrological interactions across subbasins. Besides, while the inter-subbasin
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53 643 coupling operates at a consistent subbasin scale through the same or different channel
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56 644 routing algorithms, this study did not explicitly address the coupling of processes
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59 645 across nested spatial and temporal scales within each subbasin and the transfer of
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1 646 uncertainties between them, which remains central challenges in hydrology
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3 647 community. The proposed approach and the implemented modeling framework are
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6 648 extensible to support such research in the future.
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9 649 Third, while the proposed approach enables flexible configuration of model
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11 650 structure across subbasins, it currently lacks intelligent methods and tools to reduce
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14 651 the substantial modeling burdens on users (particularly non-expert users), when
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17 652 constructing appropriate model structures for specific application contexts. Future
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21 653 efforts should focus on developing domain knowledge-driven methods that
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24 654 incorporate various knowledge types, such as: (1) integration knowledge of SEIMS
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27 655 modules derived from module metadata; (2) rule-based relationships linking
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30 656 application contexts (e.g., watershed topographic characteristics and available data) to
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33 657 suitable simulation algorithms; and (3) case-based reasoning to determine appropriate
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36 658 simulation units, algorithms, and parameter settings (Jiang et al., 2019; Qin et al.,
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38 659 2025).
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41 660 Fourth, the current case study was designed as a simplified proof-of-concept
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44 661 application to demonstrate the methodological feasibility of the proposed approach.
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47 662 Nevertheless, more comprehensive case studies in watershed with multi-scale
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50 663 observed data and diverse climatic, topographic, and landuse conditions are needed to
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53 664 further evaluate the robustness, generalizability, and uncertainty implications of the
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56 665 proposed approach. For instance, future applications should incorporate uncertainty
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59 666 analysis frameworks to better understand and quantify the trade-offs between
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1 667 modeling flexibility and simulation accuracy.
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4 668 **5 Conclusion**

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7 669 This study proposed a new spatially hybrid hydrological modeling approach to
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10 670 address the challenge of representing spatial heterogeneity in complex hydrological
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13 671 modeling application contexts. Adopting a divide-and-conquer strategy, the approach
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16 672 allows for flexible combination of compatible spatial discretization schemes (i.e.,
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19 673 simulation units) and hydrological process representations (i.e., simulation
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22 674 algorithms) across subbasins. This approach provides a pragmatic and flexible
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25 675 modeling solution to enable the integration of existing hydrological modeling
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28 676 methods at the subbasin scale.

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31 677 The approach was implemented within the SEIMS framework, leveraging its
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34 678 modular, subbasin-independent architecture and parallel computing capabilities.
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37 679 Comparative experiments in a medium-sized watershed demonstrated its integration
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40 680 feasibility and simulation effectiveness. The spatially hybrid model structure
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43 681 successfully combines the strengths of both spatially consistent model structures,
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46 682 improving general performance, while also inheriting some limitations from its
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49 683 constituent components. The experimental results highlight its capability to balance
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52 684 modeling flexibility and simulation accuracy under heterogeneous watershed
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55 685 conditions.

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58 686 This study does not intend to resolve the fundamental hydrological issues such as
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1 687 the scaling issues (Dooge, 1986; Beven, 2000), nor to ignore their existence, but
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4 688 rather to provide a practical and extensible framework for hydrological modeling
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6 689 across diverse application contexts. The proposed approach and its implementation
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9 690 are extensible through the reuse and integration of existing model efforts, thus holding
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12 691 the potential to support investigations into these theoretical issues as new theories and
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15 692 data become available. Moreover, this study lays the groundwork for future research
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18 693 on knowledge-driven intelligent modeling methods and tools for complex watershed
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15 712 **Reference**

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